



移动阅读

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湖相致密储层有利层段的沉积与湖平面变化的耦合关系及其控制因素：以柴西尕斯地区上干柴沟组下段为例

郑永盛¹, 唐闻强^{1,2}, 伊海生^{1*}, 裴梓薇³, 杨梅³, 邢浩婷³, 杨芸³, 汪素凤¹

(1. 成都理工大学沉积地质研究院, 四川 成都 610059; 2. 中国石油西南油气田分公司, 四川 成都 610051; 3. 中国石油青海油田分公司, 甘肃 敦煌 736202)

摘要: 探讨有利层段的沉积与湖平面变化之间的关系, 可以对湖相致密储层的勘探开发部署提供一定参考。以柴达木盆地西部尕斯地区上干柴沟组下段为例, 利用自然伽马 (GR) 曲线资料, 进行 Fischer 图解和铀含量 (U) - 有机碳 (TOC) 回归拟合法, 重建渐新世沉积时期湖平面变化过程, 结合磁化率资料, 讨论了湖平面变化的气候控制作用。结果表明: (1) 尕斯地区上干柴沟组下段泥质含量和 TOC 含量变化趋势一致, 均反映其在沉积过程中经历了两次湖退和湖进过程, 有利于烃源岩和储层的形成; (2) 湖平面上升期形成的高水位沉积体系, 对应于致密砂岩类储层中优质烃源岩的富集层段, 能有效形成源-储共生配置; (3) 渐新世湖平面变化受西风条件下的气候的控制, 西风带来的水汽, 形成相对湿润的气候条件, 引起湖平面的上升, 有利于烃源岩的形成, 且与有利层段的沉积呈现耦合关系。

关键词: 柴达木盆地; 湖相致密储层; 湖平面变化; 有利层段; Fischer 图解; 自然伽马

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Coupling relationship between sedimentation of favorable intervals and lake level change and its controlling factors in lacustrine tight reservoir: A case study of the Lower Shangganhaigou Formation in the Gasi area, western Qaidam Basin, China

ZHENG Yongsheng¹, TANG Wenqiang^{1,2}, YI Haisheng^{1*}, PEI Ziwei³, YANG Mei³, XING Haoting³, YANG Yun³, WANG Sufeng¹

(1. Institute of Sedimentary Geology, Chengdu University of Technology, Chengdu 610059, China; 2. Petrochina Southwest Oil and Gas Field Company, Chengdu 610051, China; 3. Qinghai Oilfield Company, PetroChina, Dunhuang 736202, China)

Abstract: Exploring the relationship between favorable sedimentary intervals and the change of lake level can provide some reference for the exploration and development deployment of the lake-phase dense reservoir. In this study, taking the Oligocene Lower Shangganhaigou Formation in the Gasi area of the western Qaidam Basin as an example, Fischer diagrams and U-TOC

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作者简介: 郑永盛 (1997—), 男, 硕士, 主要从事古生物与古环境研究工作。E-mail: zys19970915@163.com

通讯作者: 伊海生 (1959—), 男, 教授, 博士生导师, 主要从事沉积学研究及教学工作。E-mail: yhs@cdut.edu.cn

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regression fitting method were used to reconstruct the process of lake level change during the Oligocene sedimentary period using Natural gamma ray (GR) curve data, and the climate controlling factors of lake level change were discussed in combination with magnetic susceptibility data. The results show that: (1) The Fischer diagram of Lower Shangganachaigou Formation in the Gasi area is consistent with the trends of mud content and TOC content, all reflecting that it has undergone two lake retreat and lake advance processes during deposition, which are favorable for the formation of source rocks and reservoirs; (2) The high water level sedimentary system formed during the lacustrine rising period corresponds to the enrichment intervals of high-quality source rocks in tight sandstone reservoirs, which can effectively form a source-reservoir symbiotic configuration; (3) The change of lake level in the Oligocene was controlled by the climate under the westerly wind condition. The water vapor brought by the westerly wind formed relatively humid climate conditions and caused the rise of lake level, which was favorable for the formation of source rocks and presented a coupling relationship with the deposition of favorable intervals.

Key words: Qaidam basin; lacustrine tight reservoir; lake level change; favorable interval; Fischer plots; natural gamma-ray

0 引言

自 2006 年北美威利斯顿盆地 Bakken 组和德克萨斯 Eagle Ford 致密油勘探取得重大突破以来,致密油储层逐渐为国内所认识和接受(Dow W, 1974; Williams J A, 1974; 林森虎等, 2011)。相对而言,国内的致密油勘探开发工作起步较晚,初步勘探结果表明,中国致密油资源丰富,具有良好的发展前景(杜金虎等, 2014; 徐程, 2018)。湖相致密储层作为重要的勘探目标,前人对其进行大量研究,将其分为碳酸盐岩和细粒碎屑岩两大类,其中细粒碎屑岩层系可以是烃源岩与薄层粉、细砂岩互层或紧邻的岩性组合,也可以是源—储紧密接触的致密层段(贾承造等, 2012; 马洪等, 2014),考虑到源—储共生关系,烃类排出后仅需近距离运移,就能直接聚集在邻近的致密储层中(庞正炼等, 2012; 邹才能等, 2012; 董国栋等, 2013),因此,湖相致密油气的源—储空间叠置关系与控制因素的研究,成为人们关注的热点。其中,湖平面变化对源—储空间的叠置具有重要的控制作用。前人研究表明,湖平面的变化控制着湖泊氧化还原界面的升降,从而影响着烃源岩和致密储层的发育(Bao et al., 2017; 金强等, 2000; 侯启军等, 2009; 张葳, 2013; 徐延康, 2014; 李禹成, 2019; 陈云等, 2021),而湖平面变化可能会受到多个因素影响,如海侵事件(冯子辉等, 2009, 2015)、降雨(李禹成, 2019; 杨大明, 2017)、区域构造活动(郑敬贵等, 2014; 吴伟等, 2012)、古气候控制(吴伟等, 2012)、缺氧事件(韩刚等, 2012; 曹怀仁, 2017)等。因此,探讨湖相致密储层对湖平面变化的响应和湖平面变化的控制因素,有利于致

密油勘探开发的部署工作,具有重要的理论意义和实践价值。沉积特征与沉积环境分析是研究湖平面变化过程的传统方法(Harrison et al., 1993; Bookman et al., 2004; Wang et al., 2012; 王亚青, 2009),岩性变化、地球化学参数(黏土矿物、同位素和元素含量及比值)、古生物(介形类、植物化石和孢粉)、磁化率(MS)、总有机碳(TOC)、测井资料、地震资料等信息作为替代指标被广泛应用于重建湖平面变化过程(Bellanca et al., 1992; Harrison et al., 1993; Hannon et al., 1997; Leng et al., 2004; Xi et al., 2011; Liu, et al., 2013; 施之新, 1997; 何胡军等, 2003; 伊海生等, 2006, 2009; 吴伟等, 2011; 孙晶等, 2012; 林孝先等, 2014; 郑敬贵等, 2014; 李启来, 2017)。

柴达木盆地作为中国西部重要的含油气盆地,连续沉积了巨厚层的新生代地层,蕴藏着丰富的油气资源(付锁堂等, 2013, 2016; 刘占国等, 2017)。近年来,该盆地的致密油气勘探开发得到广泛重视(雷群等, 2008; 付锁堂等, 2013, 2016; 刘顺宇和赵荣, 2019)。虽然总体仍处于初期阶段,但对致密储层,特别是湖相致密储层的勘探已取得了突破性进展(Rashid et al., 2015; Xu et al., 2017; Kong et al., 2018; Liu and Xiong, 2021; 黄第藩等, 2003; 付锁堂等, 2013, 2016; 邹才能等, 2010, 2013, 2018; 马达德等, 2019),已有证据表明,自古近纪以来,柴达木盆地一直处于封闭的咸化湖盆环境,有着相对稳定的陆源输入,发育有砂岩和碳酸盐岩两类致密储层,致密碳酸盐岩储层主要分布在柴西部的狮子沟、南翼山和大风山等地区,致密砂岩储层主要分布在花土沟、尕斯、昆北和扎哈泉等地区(图 3a)(付锁堂等, 2013, 2016; 王艳清等, 2014; 石金华等, 2016;

刘占国等, 2017; 庞正炼等, 2018a, 2018b; 张道伟等, 2019), 其中尕斯地区上干柴沟组下段致密砂岩储层平均孔隙度介于 5.6%~8.1% 之间, 渗透率在 $0.01 \times 10^{-3} \sim 0.1 \times 10^{-3} \mu\text{m}^2$ 之间, 至今没有实现油气勘探的大规模突破(郭华粘, 2020), 因此, 还需开展进一步研究, 明确湖平面变化对源—储配置关系的控制, 为有利区预测提供可靠依据。据此, 本次研究以尕斯地区上干柴沟组下段为例, 利用测井资料和磁化率资料, 采用铀含量(U)—有机碳(TOC)回归拟合法和 Fischer 图解, 厘清湖平面变化与致密储层有利层段的关系, 探讨湖平面变化的气候控制因素, 以期为有利目标预测提供地质依据。

1 地质背景

柴达木盆地(北纬 $36^\circ \sim 39^\circ$, 东经 $90^\circ \sim 98^\circ$)是青藏高原北缘的一个菱形山间盆地, 以祁连山、昆仑山和阿尔金山为界(图 1), 盆地面积 $120\,000 \text{ km}^2$, 平均海拔 $2\,800 \text{ m}$ (Wu et al., 2021)。该盆地是在印度—欧亚大陆碰撞、青藏高原多期隆升的综合效应下形成的新生代沉积盆地(Li et al., 2014; Bian et al., 2019; Xiong et al., 2021), 盆内新生代沉积广泛、厚度巨大。尕斯断陷位于柴达木盆地西部坳陷带, 是坳陷内的次级构造单元, 也是盆地西部油气最富集的地区之一(图 1)(丁文龙等, 2004)。

柴达木盆地新生代自下而上依次发育古近系路乐河组(E_{1+2})、下干柴沟组下段(E_3^1)、下干柴沟组上段(E_3^2), 新近系上干柴沟组(N_1)、下油砂山组(N_2^1)、上油砂山组(N_2^2)、狮子沟组(N_2^3), 以及第四系的七个泉组(Q_{1+2})(图 2a)(Sun et al., 2005; Tang, et al., 2021; 杨藩等, 1992; 张伟林, 2006)。依据上干柴沟组中部普遍发育的自然伽马的高值为标志, 将其进一步划分为上干柴沟组下段(N_1^1)和上段(N_1^2)(图 2b)。尕斯断陷地层发育齐全, 下干柴沟组沉积期, 物源供给充分, 研究区沉积环境由下段的滨湖相过渡为上段的浅湖相; 上干柴沟组沉积期构造活动不强烈, 研究区发育滨湖—半深湖环境, 岩性主要为灰色砂岩—泥岩不等厚互层, 夹少量泥灰岩(图 2a、图 3a)(郭华粘, 2020); 下油砂山组早期构造活动继承了上干柴沟组时期状态, 区域构造活动依然不强烈, 至下油砂山组晚期, 构造活动逐渐增强, 盆地南部的昆仑山持续隆升, 柴西南地区随之抬升, 湖泊中心也逐渐迁移并远离柴西南地区(Molnar et al., 1993; Fang et al., 2005; 袁亚娟等, 2010; 王亚东等, 2011; 李俊武, 2016)。

2 材料和方法

2.1 Fischer 图解

Fischer 图解通过纵轴上旋回厚度的累积残差

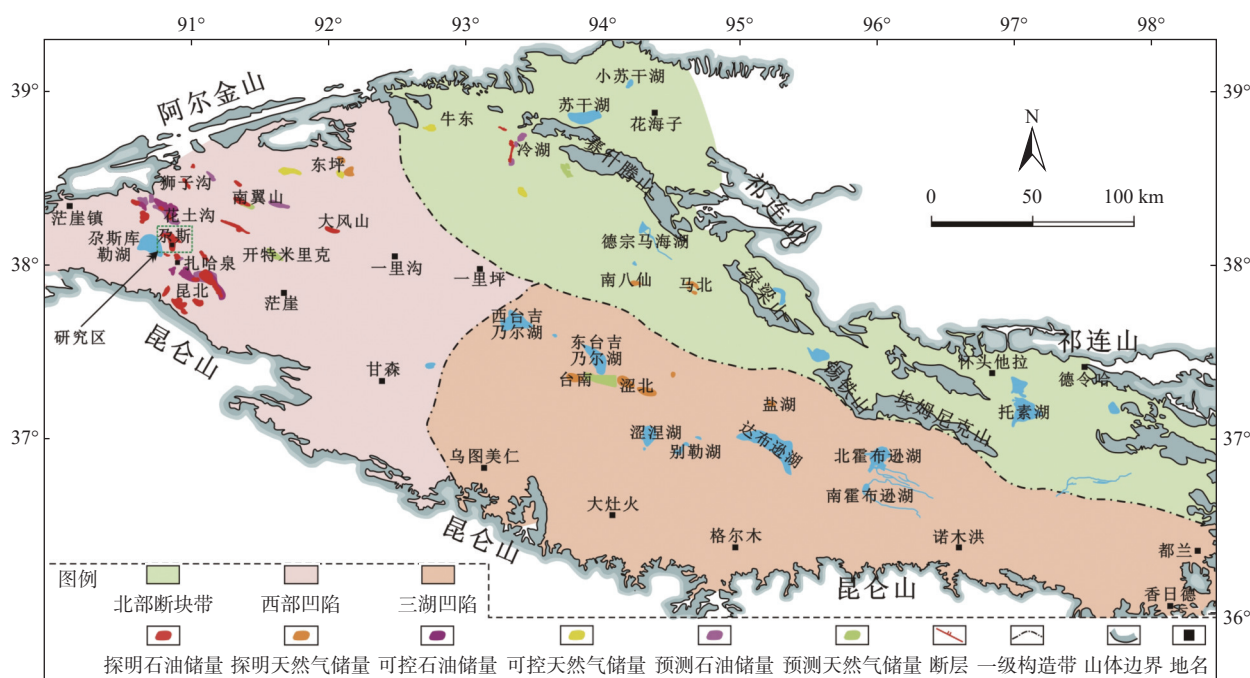


图 1 柴达木盆地构造单元 (据付锁堂等, 2013 修订)
Fig. 1 Tectonic units of the Qaidam Basin (modified from Fu et al., 2013)

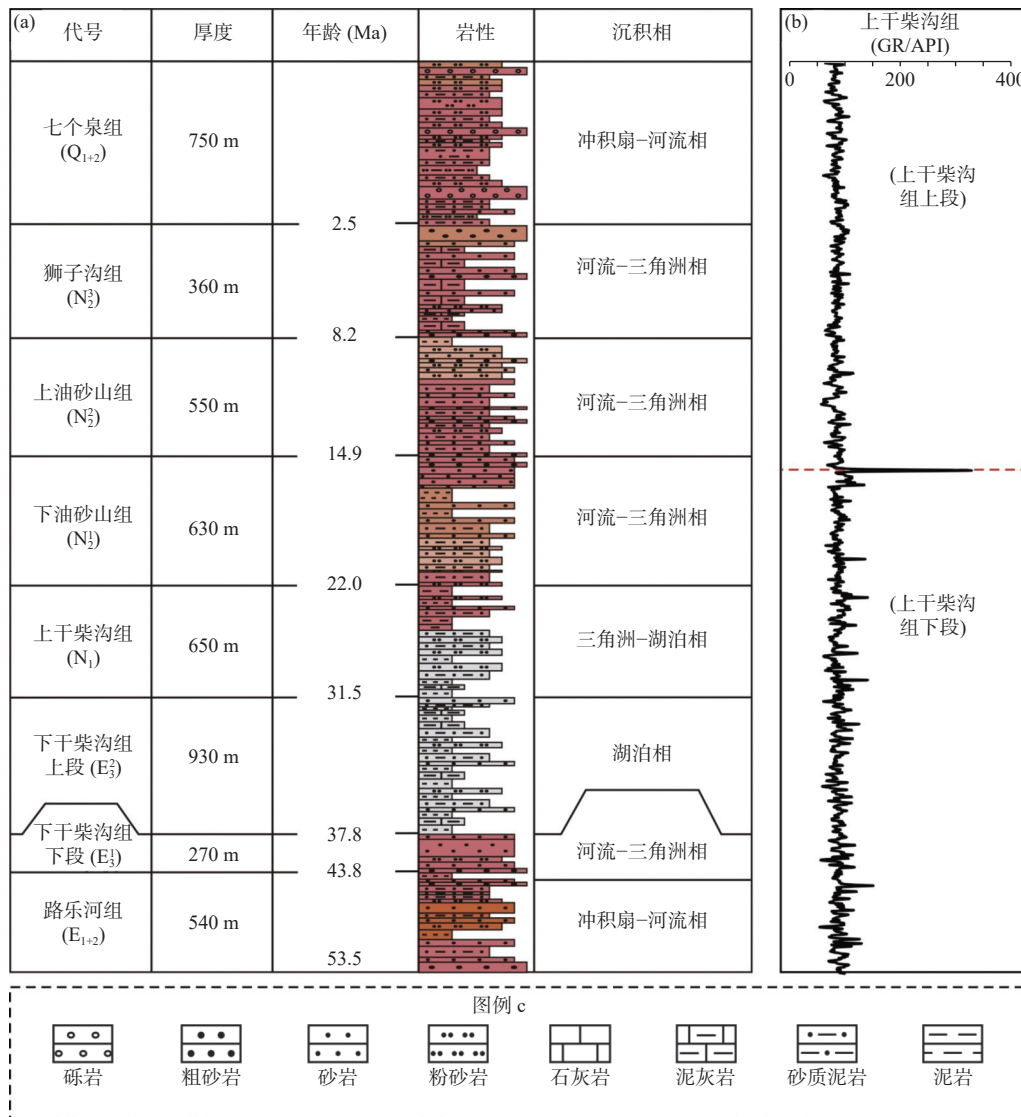


图 2 柴达木盆地西部尕斯地区地层 (a) 和上干柴沟组测井曲线 (b)

Fig. 2 Generalized stratigraphic column of the Gasi area, western Qaidam Basin(a) and GR logs in the LSG Formation (b)

和横轴上旋回数(或深度域)构建,代表沉积物形成时有效可容纳空间的变化,故 Fischer 图解又称为可容纳空间图解(Fischer, 1964; Read and Goldhammer, 1988; Wang et al., 2009; 梅冥相等, 2001; 武向峰, 2010; 龚大兴等, 2014)。在湖相沉积体系中,可用自然伽马作为替代指标(Yin et al., 2015; Yang et al., 2018; 伊海生, 2011)表征可容空间的变化,进而反映湖平面的波动。本次研究基于研究区 9 口探井的测井资料,首先对原始 GR 资料进行局部回归平滑(LOWESS 和 LOESS)处理,去除测井曲线的长期趋势和降噪;然后利用一阶差分法使数据中心化或归一化,保证最终计算结果具有一致性;再通过逻辑判别函数统计旋回厚度与个数;最后利用层段内平均旋回厚度,计算各旋回的厚度

偏差,然后通过累加厚度偏差形成 Fischer 图解(图 4)。测井资料来自中国石油天然气集团公司的 EI-Log 测井系列,间距大多为 0.1 m,少数为 0.125 m。

2.2 磁化率 (MS)

磁化率(MS)是表征物质在外加磁场作用下的磁化强度(Shi et al., 2018; Zhang et al., 2019),大尺度的变化趋势可能反映了气候变化(Bao et al., 2018)。在相对湿润的环境中形成具有高磁化率和细颗粒尺寸的沉积物,表明水位较高;反之,在相对干旱的环境中,形成具有较低磁化率和较粗粒度的沉积物,表明水位较低(胡守云等, 1998)。本次研究中,使用手持式磁化率仪 KT-10,灵敏度为 10^{-6} SI,以间隔 2 m 的采样间距,重点对 Y9 井上干柴沟组下段岩屑样品进行磁化率的实测,共获得 158 个

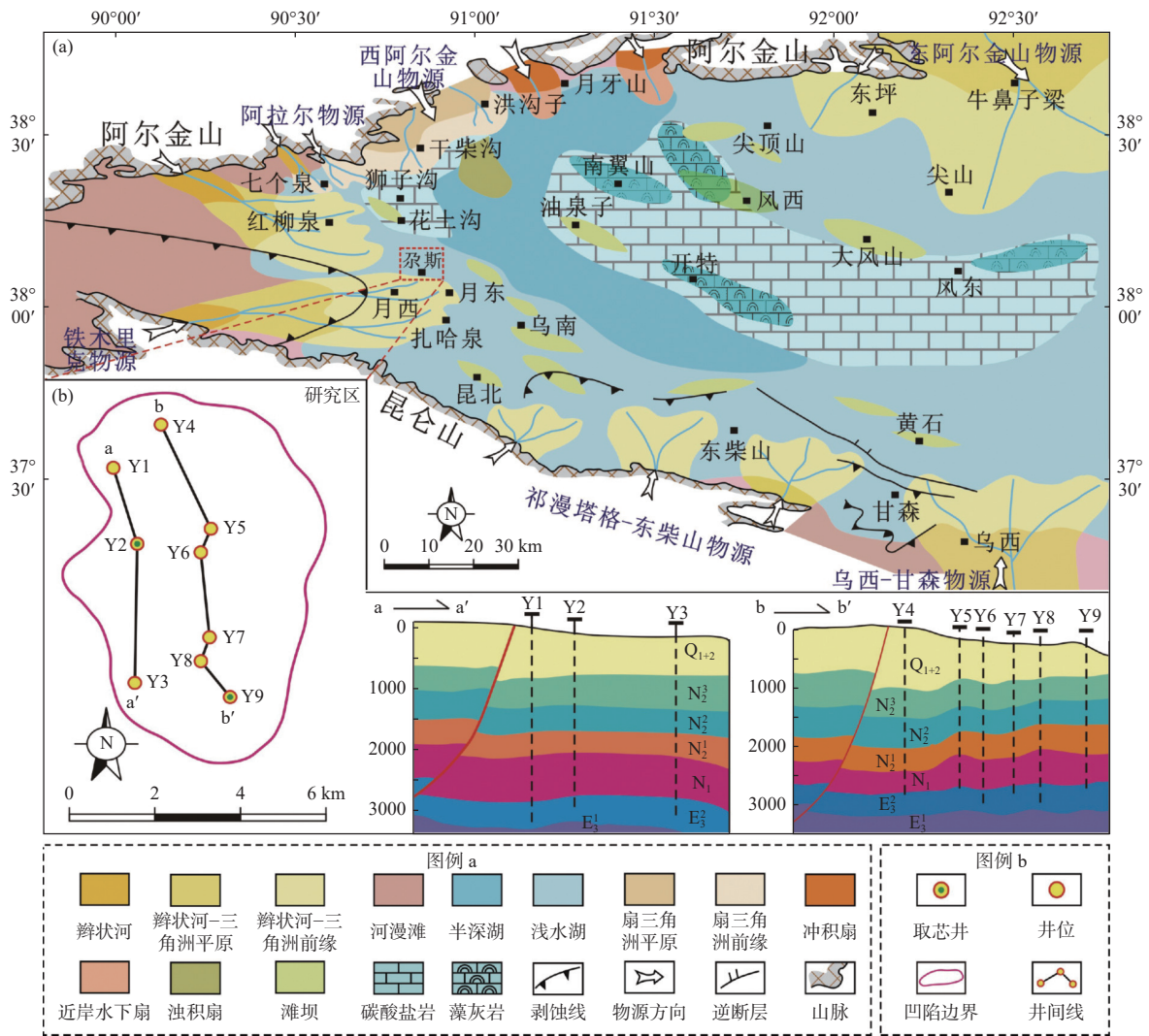


图3 柴达木盆地西部上干柴沟组沉积体系图(a, 据王艳清等, 2014 修改)和连井剖面及相应地震剖面示意图(b)
Fig. 3 Map of the sedimentary system of the Shanggancaigou Formation in the western Qaidam Basin (a; modified from Wang et al., 2014) and location map of cross-well lines and schematic diagram of the corresponding seismic profile (b)

磁化率值。

2.3 铀含量 (U) —有机碳 (TOC) 回归拟合公式

沉积岩中有机碳 (TOC) 含量变化是致密油烃源岩重要的评价指标, 其受湖平面变化的控制 (王艳清等, 2014; 郑茜等, 2015)。郑茜等 (2015) 在研究柴达木盆地与尕斯地区相邻的扎哈泉地区新上干柴沟组烃源岩评价时, 利用测井资料分析了铀含量 (U) 与有机碳 (TOC) 之间存在良好的指数关系, 并建立了 U—TOC 回归拟合经验公式:

$$TOC = 0.4383 \times e^{0.9707\Delta U} \quad (1)$$

$$\Delta U = \frac{U - U_{\min}}{U_{\max} - U_{\min}} \quad (2)$$

尕斯地区与扎哈泉地区邻近, 均受到盆地西部物源影响 (图 3a) (王艳清等, 2014)。为更好地反映

湖平面变化与烃源岩发育层段的相互关系, 研究采用该 U—TOC 经验公式, 计算尕斯地区 TOC 含量的变化趋势, 为致密油烃源岩的评价提供可靠的依据。

公式中铀含量 (U) 是利用自然伽马能谱测井获得, 测井仪器为哈利伯顿的 LOG-IQ 系列, 采样间距为 0.1 m。

3 结果

3.1 Fischer 图解重建湖平面

基于识别的高频沉积旋回, 获得的累积旋回厚度偏离的最终曲线是深度域中的 Fischer 图解。研究中首先利用尕斯地区标准井 Y1 井, 对上干柴沟组下段进行了 Fischer 图解分析 (图 4), 结果显示,

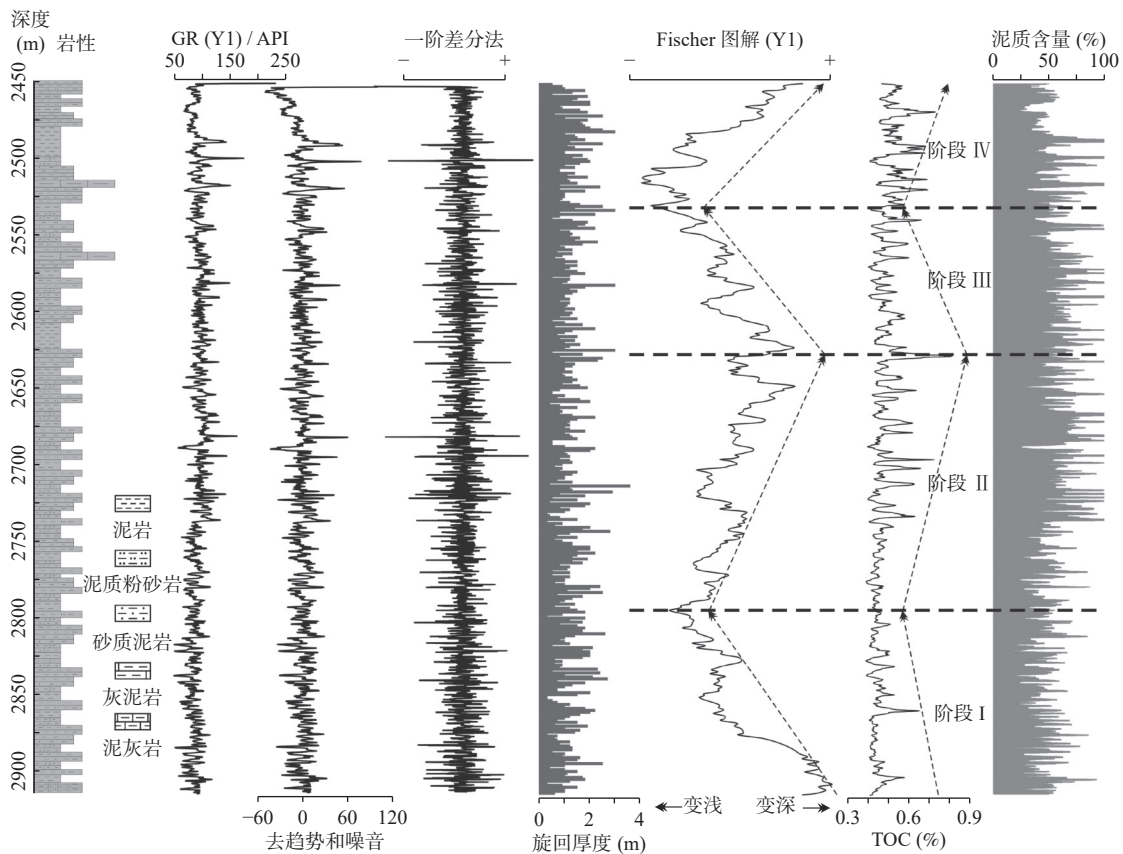


图 4 Y1 井上干柴沟组下段的 Fischer 图解、TOC 和泥质含量对比 (20% 加权平均值用于去趋势)

Fig. 4 Comparison of Fischer plot, TOC, and shale content in the LSG Formation of well Y1. The 20% weighted average was used for detrending

尕斯地区上干柴沟组下段在沉积过程中经历了两期湖平面升降旋回, 并且湖平面的升降变化与地层中泥质含量的变化之间存在良好的对应关系(图 4), 亦即: (1) 第一旋回为 2 915.8 m~2 628.9 m, 其中 2 915.8~2 795.7 m 为湖平面下降期, Fischer 图解显示湖平面逐渐变浅, 对应的泥质含量也呈由高到低的变化特征; 2 795.7~2 628.9 m 为湖平面上升期, Fischer 图解显示湖平面逐渐变深, 对应的泥质含量也呈由低到高的变化特征。(2) 第二旋回为 2 628.9~2 453.3 m, 其中 2 628.9~2 532.8 m 为湖平面下降期, 同样的 Fischer 图解显示湖平面逐渐变浅, 对应的泥质含量也呈由高到低的变化特征; 2 532.8~2 453.3 m 为湖平面上升期, Fischer 图解显示湖平面逐渐变深, 对应的泥质含量也呈由低到高的变化特征(图 4)。

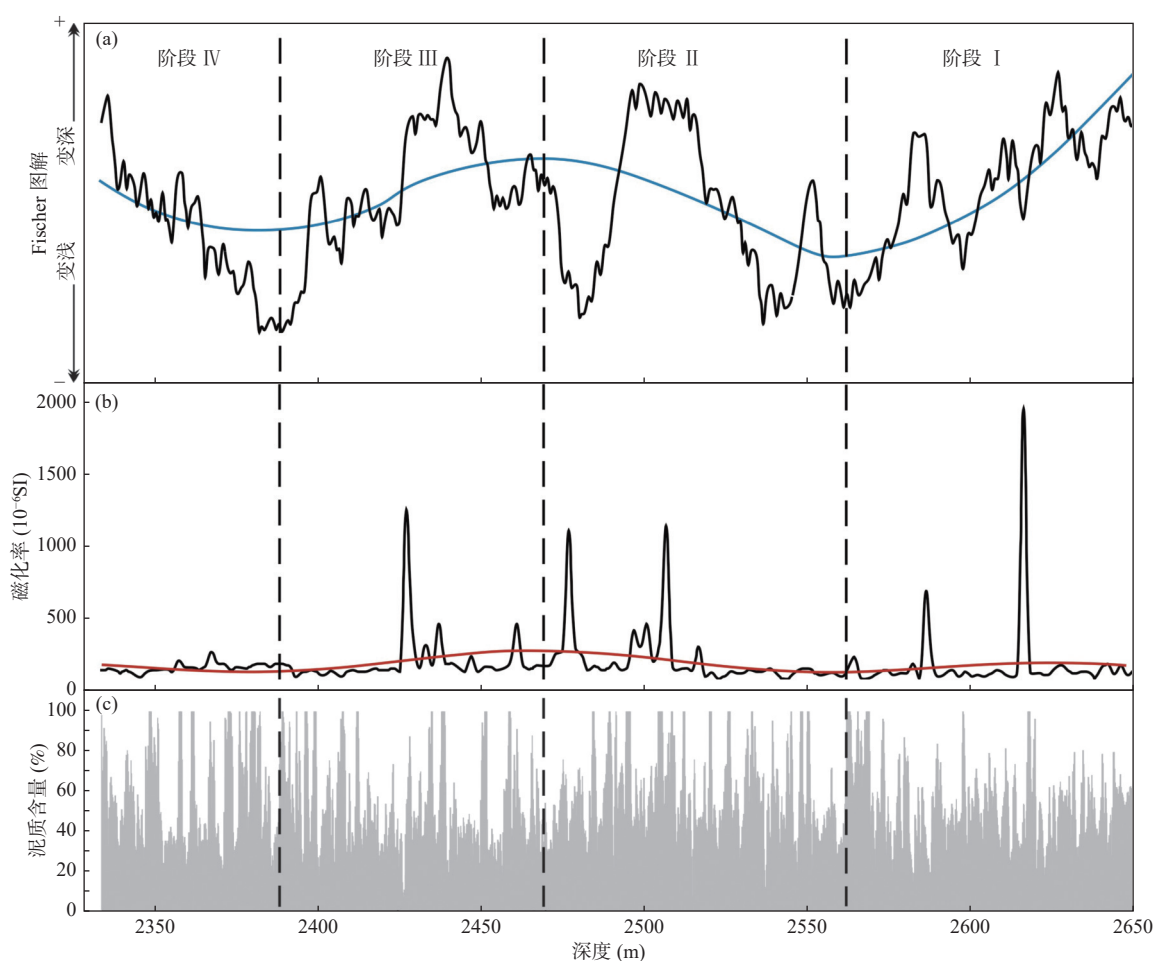
3.2 磁化率 (MS)

尕斯地区 Y9 井岩屑样品的磁化率实测数据如图 5 所示, 依据 Fischer 图解划分为四个阶段。结果显示, 磁化率与 Fischer 图解显示湖平面变化

和泥质含量的变化之间存在良好的对应关系(图 5)。第一阶段(2 559.7~2 650 m)磁化率范围在 80.8×10^{-6} ~ 1969×10^{-6} SI, 平均值为 176.63×10^{-6} SI, 呈现下降趋势, 对应于湖平面下降与泥质含量降低; 第二阶段(2 482.9~2 559.7 m)磁化率范围为 79.3×10^{-6} ~ $1 145.9 \times 10^{-6}$ SI, 平均值为 188.25×10^{-6} SI, 呈现上升趋势, 湖平面处于上升期, 泥质含量升高; 第三阶段(2 387.7~2 482.9 m)磁化率范围在 91.8×10^{-6} ~ $1 255.4 \times 10^{-6}$ SI, 平均值为 222.28×10^{-6} SI, 对应于湖平面第二次下降期, 泥质含量由深到浅; 第四阶段(2 332.0~2 387.7 m)磁化率范围在 90.8×10^{-6} ~ 270.1×10^{-6} SI, 平均值为 156.08×10^{-6} SI, 呈现上升趋势, 与湖平面变化和泥质含量变化趋势一致(图 5)。

3.3 TOC 含量的计算

根据经验公式, 计算出 Y1 井上干柴沟组下段的 TOC 呈现明显的分段性(图 4), 并且从图中可以看出其与湖平面的升降变化与地层中泥质含量的变化之间存在显著的正相关关系。第一阶段(2 795.7~2 915.8 m), TOC 在 0.39%~0.65% 之间变



蓝线表示 Fischer 图解加权平均值的去趋势；红线表示磁化率 (MS) 加权平均值的去趋势。

图5 Y9井上干柴沟组下段 Fischer 图解 (a)、磁化率 (MS) 趋势 (b) 与泥质含量 (c) 对比

Fig. 5 Comparison of Fischer plot (a), magnetic susceptibility (MS)(b), and shale content (c) in the LSG Formation of well Y9. The blue line shows the detrending of the weighted average of the Fischer plot. The red line shows the detrending of the weighted average of the MS

化, 平均值为 0.45%, 呈下降趋势, 对应于湖平面和泥质含量下降期; 第二阶段(2 628.9~2795.7 m), TOC 在 0.39%~0.81% 之间变化, 平均值为 0.47%, 呈上升趋势, 处于泥质含量和湖平面上升期; 第三阶段(2 532.8~2628.9 m), TOC 在 0.40%~0.80% 之间变化, 平均值为 0.48%, 与泥质含量和湖平面呈现第二次下降趋势; 第四阶段(2 453.3~2 532.8 m), TOC 在 0.41%~0.73% 之间变化, 平均值为 0.53%, 三者整体处于第二次上升趋势。

4 讨论

4.1 古湖平面演化对有利储层段的响应

已有研究表明, 湖相暗色泥岩作为陆相盆地重要的烃源岩, 只有当泥岩中的有机质达到一定丰度,

且达到一定的成熟度时才可生成油气(徐延康, 2014), 而高有机质丰度往往出现在最大湖泛面的沉积物中, 这通常与湖泊氧化还原界面的升高和陆源有机质输入的增加密切相关(Kelts, 1988; Gonçalves, 2002; Meyers, 2003)。当岩心的沉积特征表明水体从深到浅直至暴露表面时, 水体中的氧含量会随着暴露表面的接近而增加, 此时, 细菌呼吸作用和无机氧化作用会增强, 从而选择性地去除有机物中的不稳定成分, TOC 含量也会随之下降; 相反, TOC 含量也就增加。图 3 表明, Y1 井上干柴沟组下段的湖平面变化整体上呈现两次上升和下降的过程, 说明湖平面变化与 TOC 含量呈显著正相关性。此外, 地质学家在世界不同地区的湖相沉积, 如维多利亚湖和鲁克瓦湖(Talbot and

Livingstone, 1989)、Bouchet湖 (Patience et al., 1996)、若尔盖盆地(张平等, 1995)以及松辽盆地(席党鹏等, 2009; 徐延康, 2014)的研究表明湖平面变化、氢指数(HI)和总有机碳含量(TOC)之间具有良好的正相关关系。Y2井的实验结果表明, HI值、TOC和生烃潜力(S_1+S_2)与Fischer图解显示的湖平面变化具有很好的正相关性, 即随着湖平面升高而显著增加的趋势(图6)。

综合湖平面变化与泥质含量和TOC含量变化研究成果, 可以看出, 三者变化趋势在尕斯地区在上干柴沟组下段是一致的。这种相关性也反映了湖平面的变化与有利储层段的沉积具有一定耦合关系。即湖平面上升, 泥质含量增加, 有机质相对富集, 烃源岩保存良好; 由于该层段物性较差(低孔超一特低渗储层), 烃源岩排出的烃类物质仅仅经过短距离运移, 便在邻近的致密砂岩储层中得以有效保存, 从而形成优势层段。这一结果也得到实际勘探的证实, 研究区9口井均具有良好产油量, 平均单井初期产油量在12 t/d(郭华粘, 2020)。

4.2 湖平面变化的气候控制因素

从地理位置上来说, 柴达木盆地自始新世以来就是一个四周多山的封闭湖泊盆地(Ye et al., 2020), 该盆地位于亚洲内陆中低纬度地区, 自晚始

新世以来可能一直处于干旱地带(Bosboom et al., 2011), 结合盆地古湖泊的规模和盐度的现有数据, 进一步表明柴达木古湖泊在上干柴沟组沉积时期可能是干旱区的封闭湖泊(局部降水和蒸发之间的平衡作用控制着干旱地区封闭湖泊的水位波动(Ye et al., 2020)。当气候湿润时, 降雨量大于蒸发量, 流入湖泊的径流增加, 同时携带更多的泥沙, 造成湖面上升; 当气候干燥时, 降雨量小于蒸发量, 入湖径流减少, 所携带的泥沙减少, 湖平面随之下降。这与前面讨论的结果是一致的, 表明尕斯地区在上干柴沟组下段沉积期的湖平面变化受到了气候控制的作用。

磁化率作为一个重要的环境指标, 在恢复古环境中起着重要作用(Thompson, 1975; Kukla, 1987)。当气候环境为相对湿润时, 降雨量大于蒸发量, 风化成壤作用强, 产生的赤铁矿就较多, 则磁化率值就较高(武向峰, 2010), 表明磁化率也可用于反映古气候环境。本研究选择Y9井的磁化率数据与Fischer图解和泥质含量进行比较和分析, 结果表明三者的变化趋势一致(图5), 气候条件相对湿润时, 磁化率高值与Fischer图解显示的高湖平面对应, 这说明在上干柴沟组时期, 尕斯地区的湖平面变化主要是受到气候控制的作用(图5)。

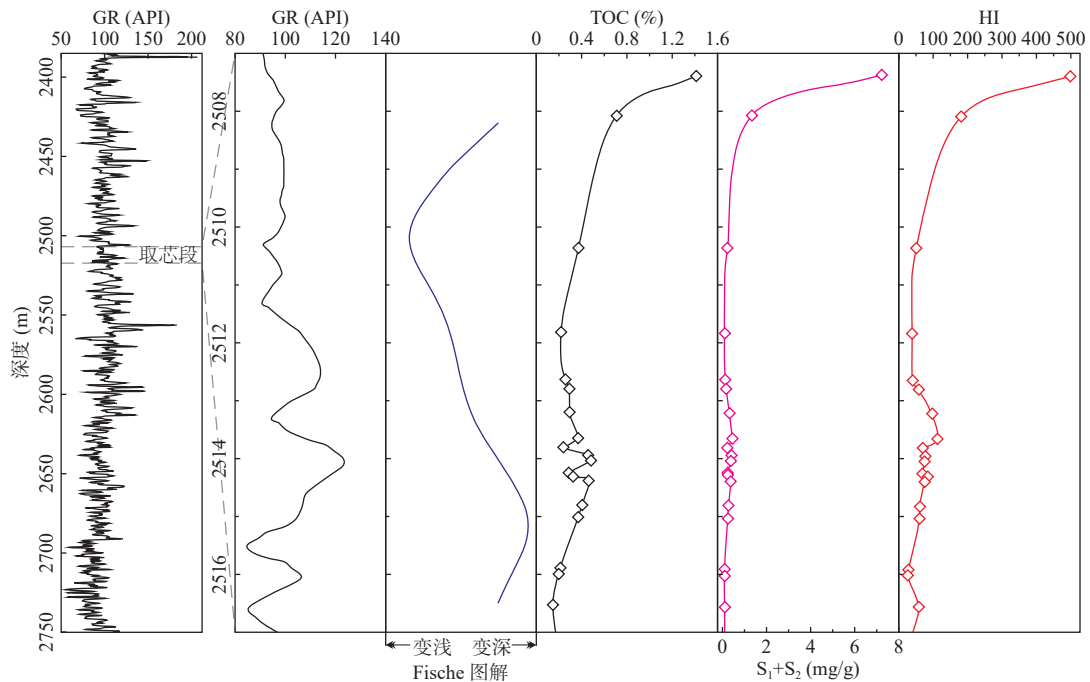


图6 Y2井岩心段Fischer图解、TOC、生烃潜力(S_1+S_2)、氢指数(HI)对比

Fig. 6 Comparison of Fischer plot, TOC, hydrocarbon generation potential (S_1+S_2), and hydrogen index (HI) in the core section of well Y2

结合已有的年代地层格架,上干柴沟组下段沉积期间处于早渐新世,该时期的深海氧同位素表明,全球因南极冰盖的扩张持续变冷(Zachos et al., 2008)(图 7a),由于南极冰盖的不断扩张,以环流和洋流的方式,造成北半球低纬地区温度降低,由赤道地区向中高纬地区输送的热量、水汽便会较少,副热带高压便会向南移动(Frierson et al., 2007; Johanson and Fu, 2009),这就增大了其与副极地低气压带的气压梯度力,从而造成西风带取代副热带高压带,为亚洲内陆带来更多的气水(图 7b)。利用叶蜡正构烷烃 $\delta^2\text{H}$ 的古水文记录,也证实了柴达木盆地西部渐新世时期的湿润气候和湖泊扩张是受西风带的影响(Wu et al., 2021)。因此,研究结果进一步表明,由于南极冰盖的扩张的影响,西风带为研究区带来了更多水汽,气候相对湿润,因而湖平面上升,有利于烃源岩的形成,湖平面变化与源—储共生配置呈现耦合关系。

5 结论

(1) 铀含量(U)—有机碳(TOC)回归拟合法计算得出的 TOC 含量变化与泥质含量变化趋势一致,上干柴沟组下段沉积期间,发生了两次湖平面上升和下降的过程,湖泊水体较为动荡,有利于烃源岩和致密储层的发育。

(2) 9 口探井的 Fischer 图解显示的湖平面变化与泥质含量变化趋势一致,且 Y1 井上干柴沟组下段的湖平面变化与 TOC 含量呈显著正相关,有利层段对应于湖平面上升导致的高水位沉积体系,该层段砂岩储层与之相邻的特性,有利于形成源—储共生配置,表明湖平面变化与致密储层有利层段的沉积具有耦合关系。

(3) Y9 井上干柴沟组下段磁化率数据与 Fischer 图解和泥质含量变化趋势一致,说明湖平面的变化与湿润气候密切相关,北半球低纬地区受南极冰盖

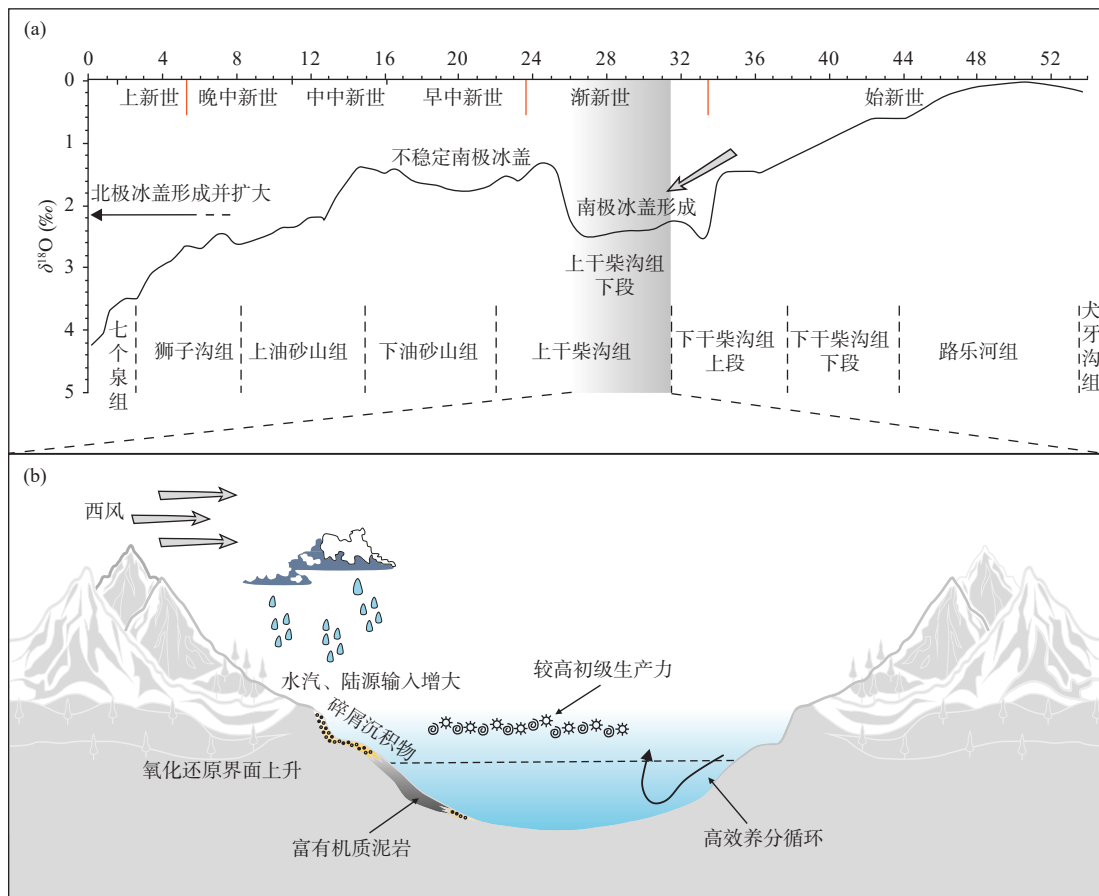


图 7 全球氧同位素含量变化 (a; 据高军平, 2009 修改) 与柴达木盆地地层划分和西风带效应降水模式图 (b)
Fig. 7 Global oxygen isotope variation (a; modified from Gao, 2009) and stratigraphic division of Qaidam Basin and westerly effect precipitation model diagram (b)

扩张影响而降温,进而导致西风带势力增强,这可能是气候变得湿润的主要原因,强劲的西风势力可为亚洲内陆盆地带来更多的水分,而水汽的增加又将利于更大降雨量的发生以及湖泊的扩张,是造成研究区湖平面变化的主要驱动力。

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